

a

Team BMO

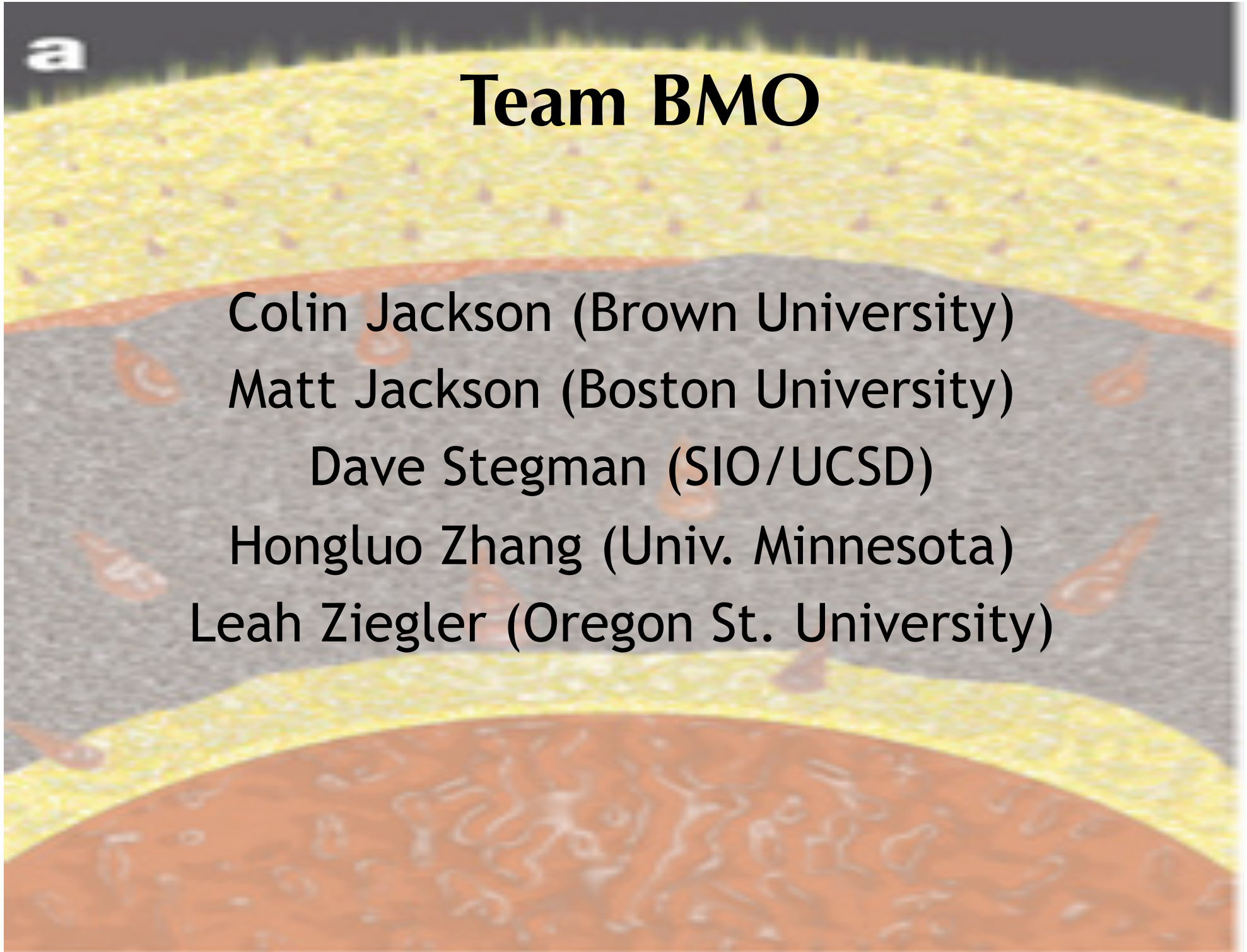
Colin Jackson (Brown University)

Matt Jackson (Boston University)

Dave Stegman (SIO/UCSD)

Hongluo Zhang (Univ. Minnesota)

Leah Ziegler (Oregon St. University)



a

Research Questions

- What is the range of plausible scenarios for solidification and stability of a BMO?
- What are the expected geochemical signatures for BMO products?
- How sensitive are these signatures to the solidification sequence?
- Can these signatures be related to observed geochemical reservoirs and/or linked to structures of the deep mantle?

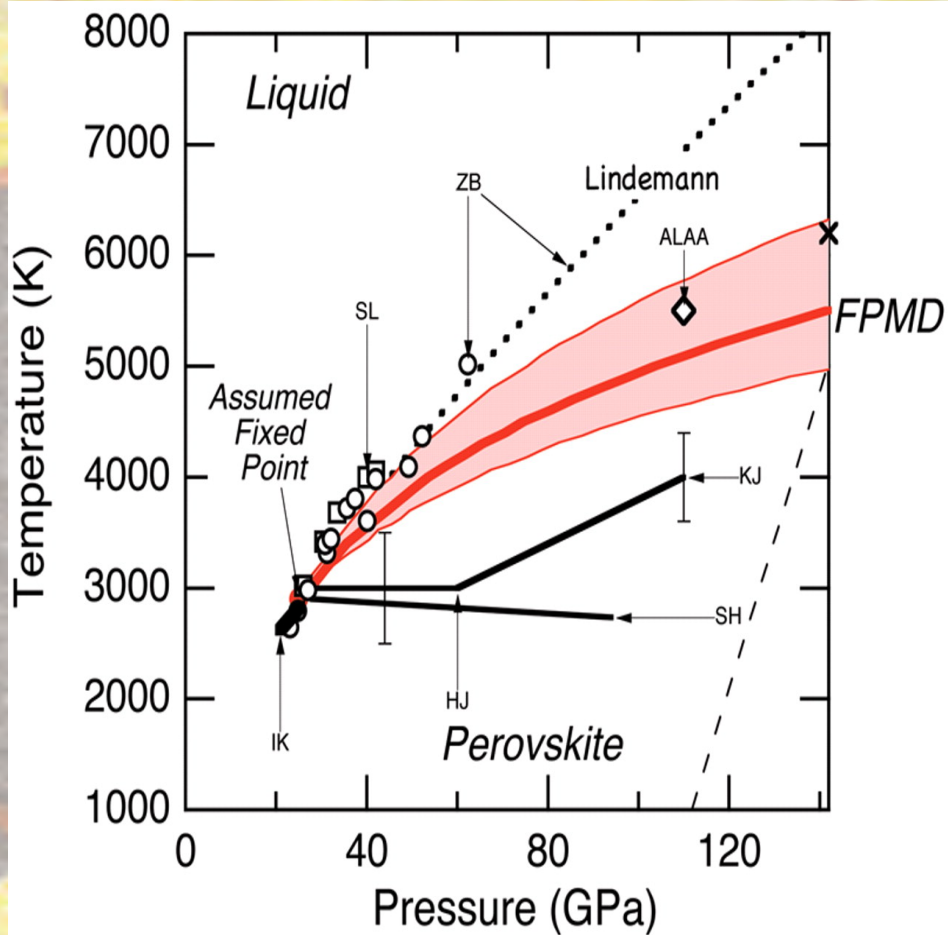
a

Past Work

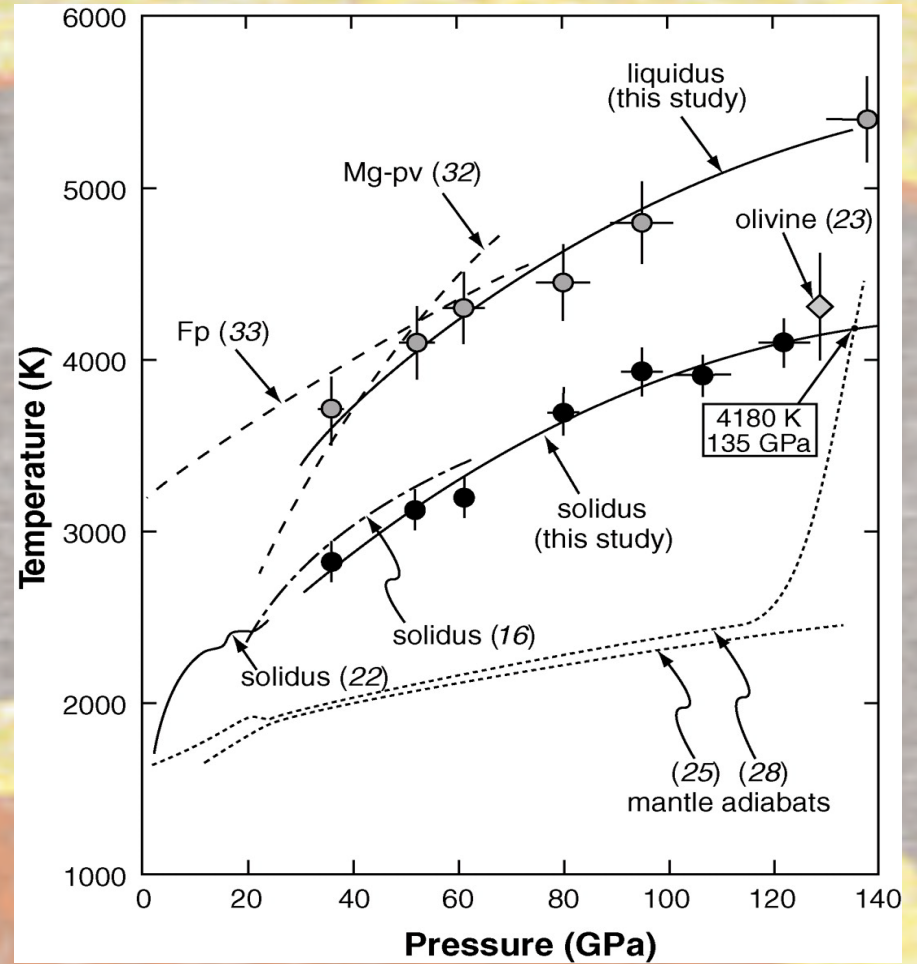
Caro et al., 2005; Labrosse et al., 2007

- Assumed constant partition coefficients
- Did not consider influence of minor elements (Ca, Al)
- Explored a very limited range of solidification scenarios (initial depth, starting composition, duration)

Estimates of Pv Liquidus

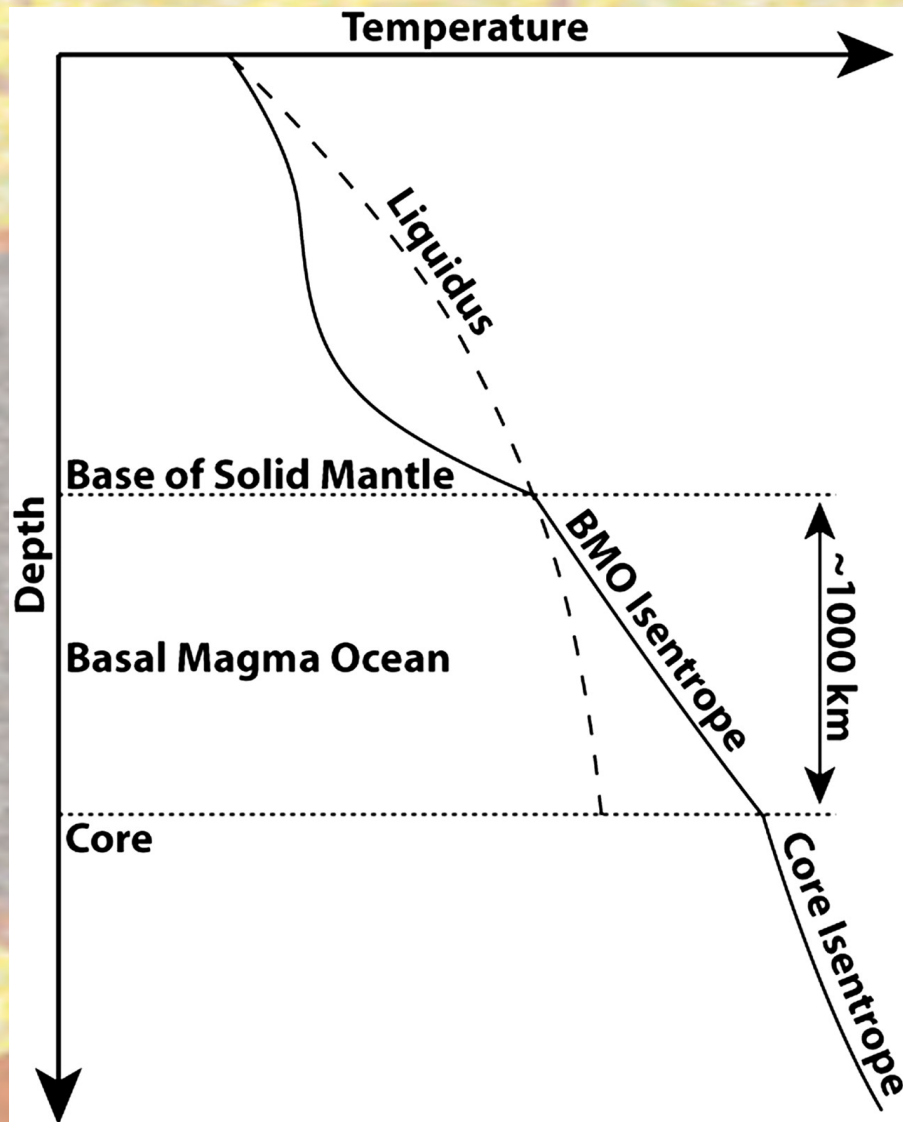


Stixrude and Karki, 2005



Fiquet et al., 2010

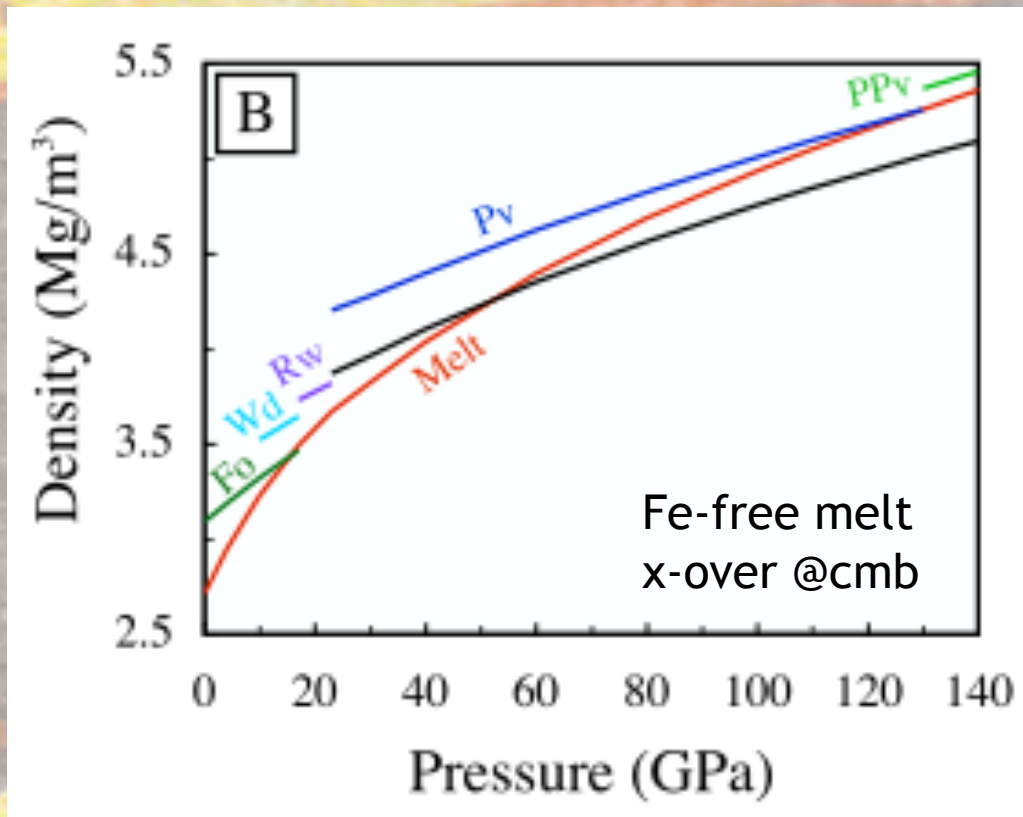
Liquidus vs Adiabats



Coltice et al, 2011

- At which P-T conditions does the liquidus intercept the adiabat of magma ocean?
- What depth does the magma ocean start crystallizing?
- Will the crystals sink or float?

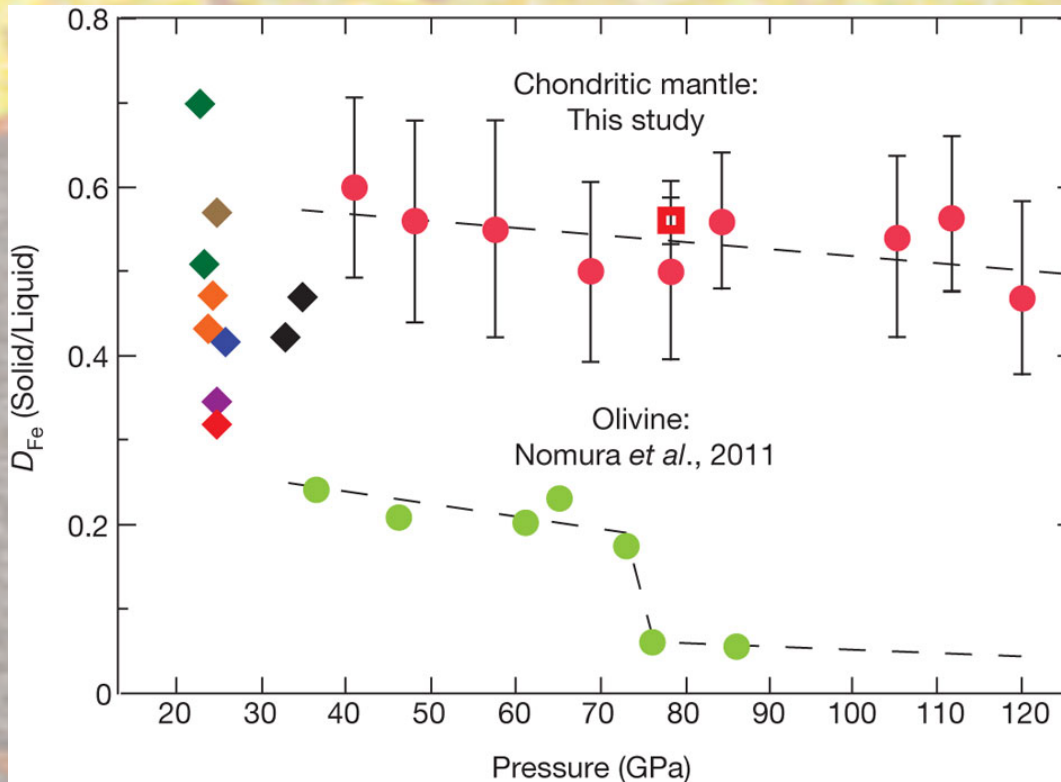
Density of melt vs solid



Mosenfelder, 2009

- How dense is the melt compared to solid?
- Added Fe to system suggests density cross-over in lower mantle
- Amount of Fe in liquid is critical to density of silicate melt and BMO depth

Behavior of Iron



- Fe-partitioning seems dependent on aluminum content
- Need to understand systematics of Al in order to predict Fe

Andrault et al, 2012

a

Research Aims of Project

Step I: Develop framework to model solidification of basal magma ocean (code named BMO2XL)

A. compile database of solid Pv-melt partition coefficients for major, minor, and trace elements

B. investigate dependence of coefficients on P-T-X; describe empirical relationships

C. apply empirical relationships to solidification sequence of basal magma ocean

Step II: Compare predicted geochemical signatures

Step III: Couple solidification history to thermal evolution model

a

BMO2XL

database:
 K_D 's for Pv

Currently has
21 references

Empirical
relationships:
 $K_D = f(P, T, X)$

Starting
BSE Comp.

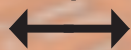
Mg/Si ?

Fractional
crystaliz.
model

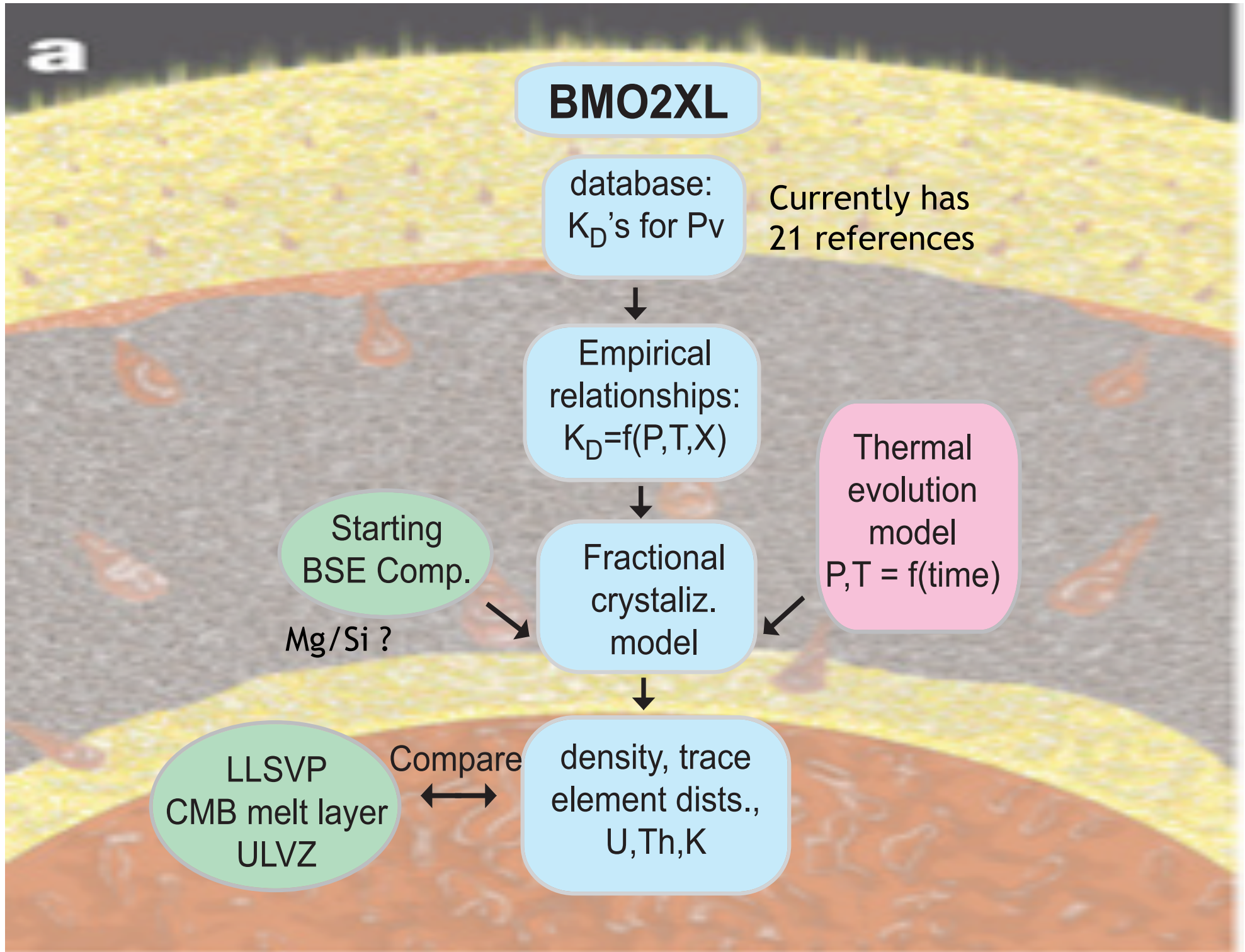
Thermal
evolution
model
 $P, T = f(\text{time})$

LLSVP
CMB melt layer
ULVZ

Compare



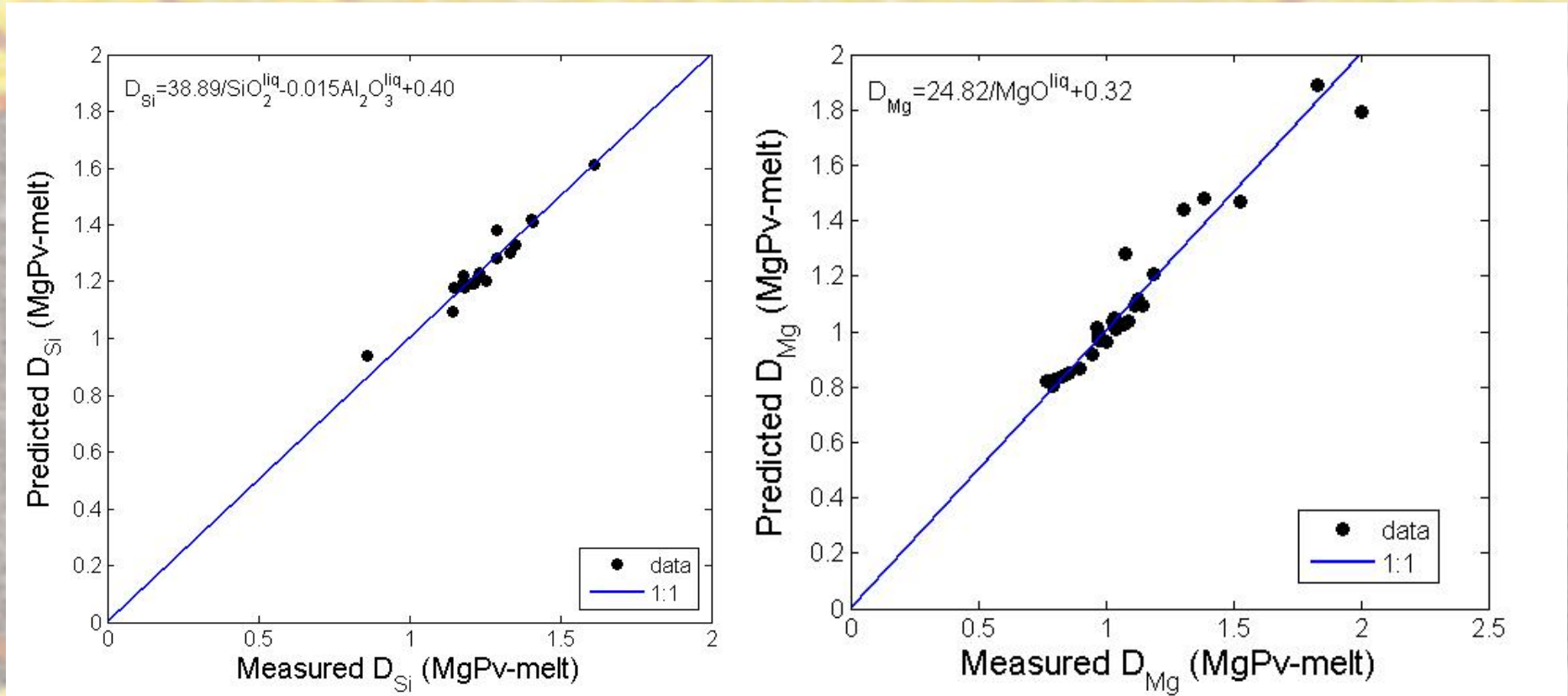
density, trace
element dists.,
U, Th, K



Results: major elements (Si, Mg)

$$D_{\text{Si}} = \text{fn}(\text{Si}, \text{Al})$$

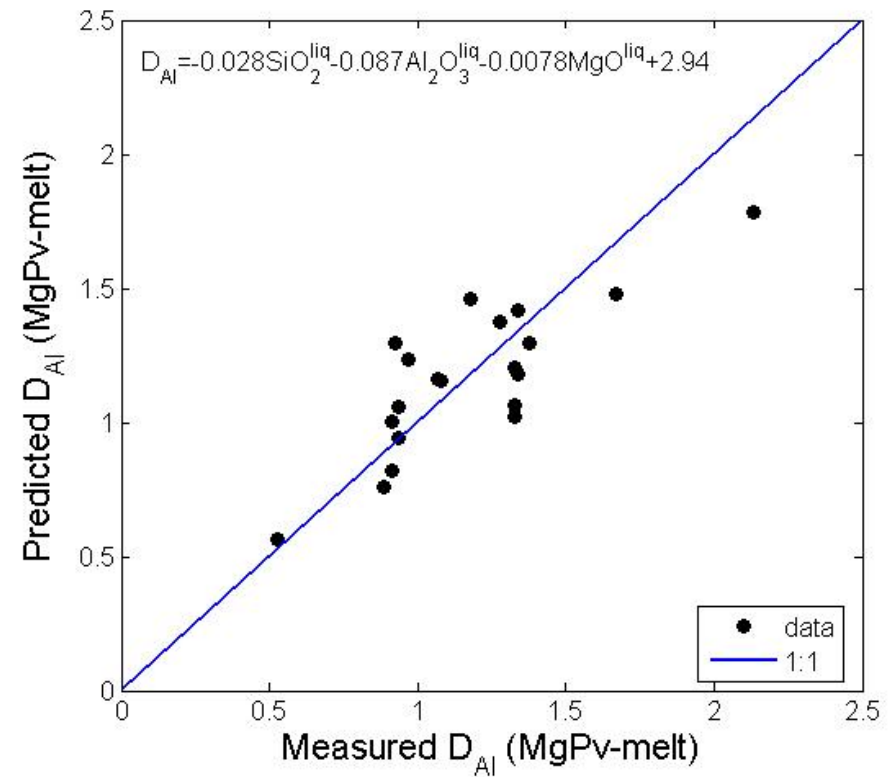
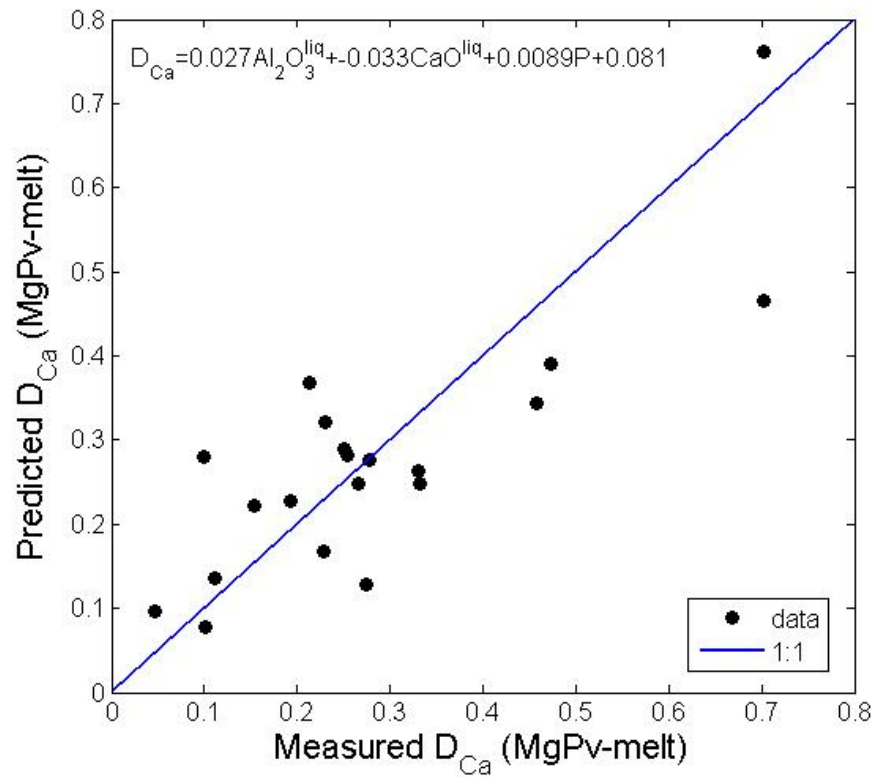
$$D_{\text{Mg}} = \text{fn}(\text{Mg})$$



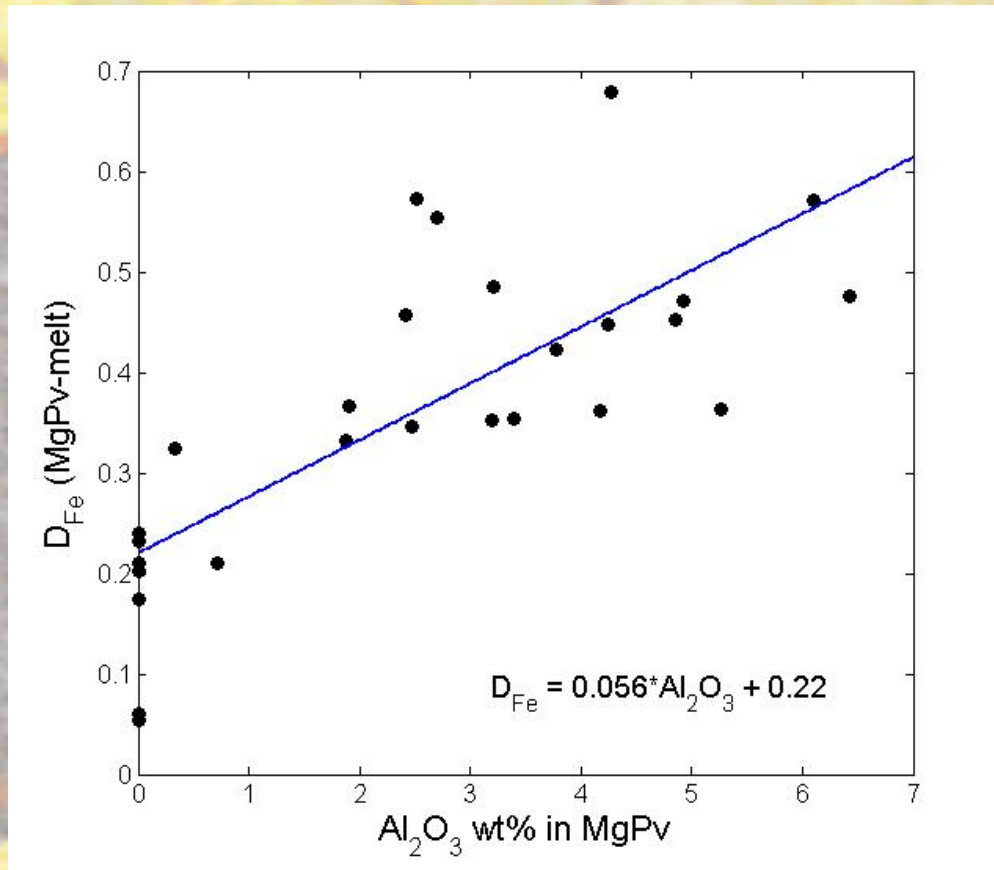
Results: major elements (Ca, Al)

$$D_{Ca} = \text{fn}(\text{Al}, \text{Ca}, \text{pressure})$$

$$D_{Al} = \text{fn}(\text{Si}, \text{Al}, \text{Mg})$$



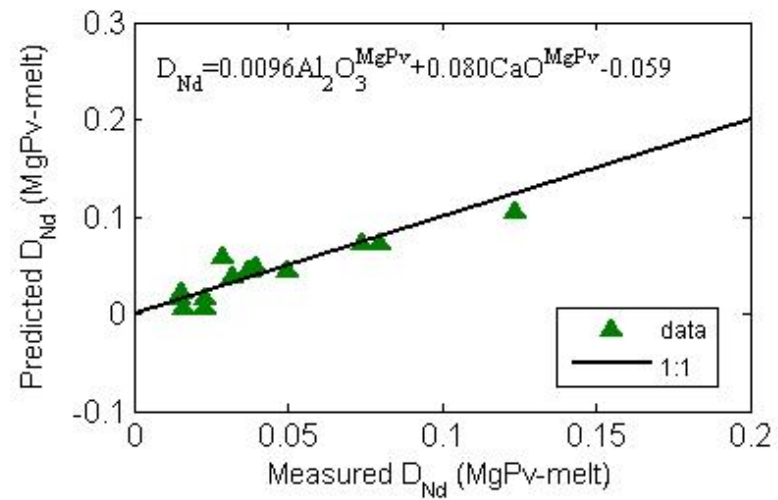
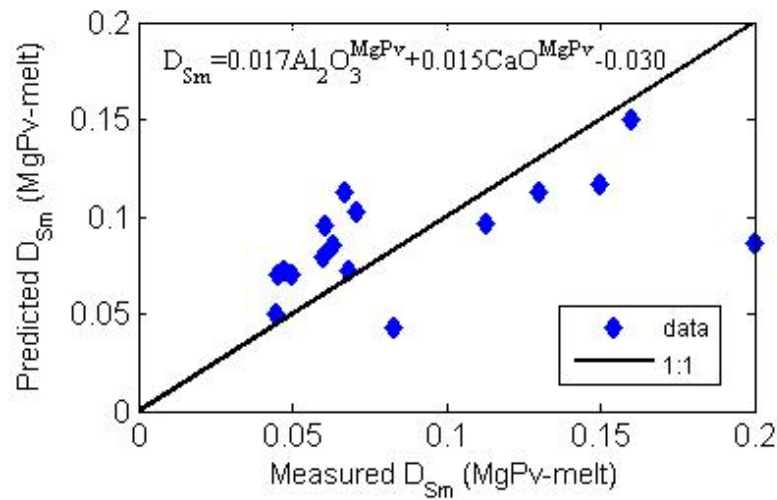
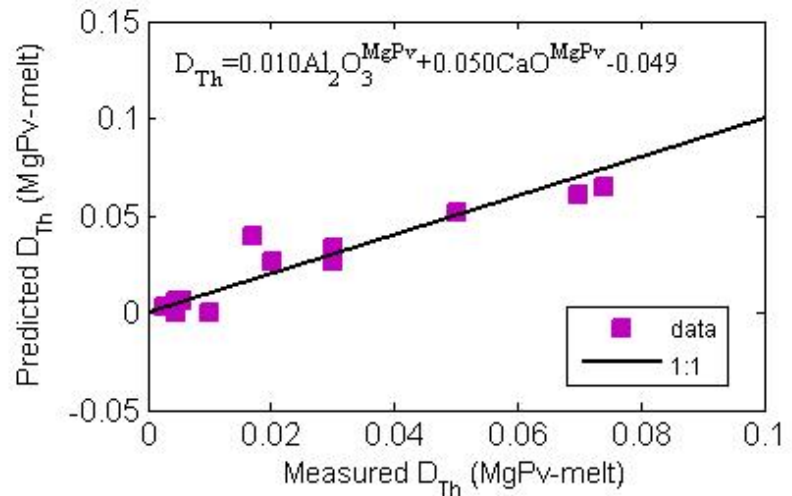
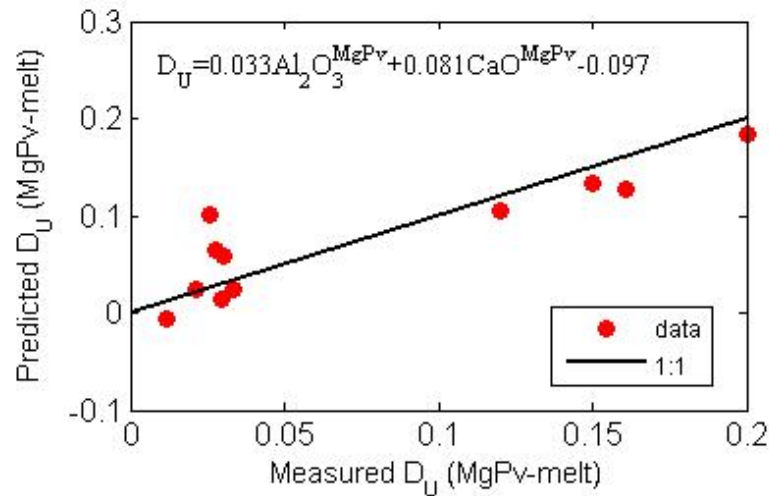
Results: behavior of Iron



- Fe-partitioning *does* depend on Al content *but* Al content depends on Si, Al, and Mg
- Fe content in melt evolves non-linearly

a

Results: trace elements



a

Conclusions

- Empirically found Al and Ca influence Fe partitioning
- Amount of Fe enrichment in silicate melt depends on solidification history
- Evolving melt composition is non-linear
- Solidification scenarios will be more complex than previous models have assumed
- Many interesting scenarios yet to be explored
- Profound implications for Earth's thermal evolution